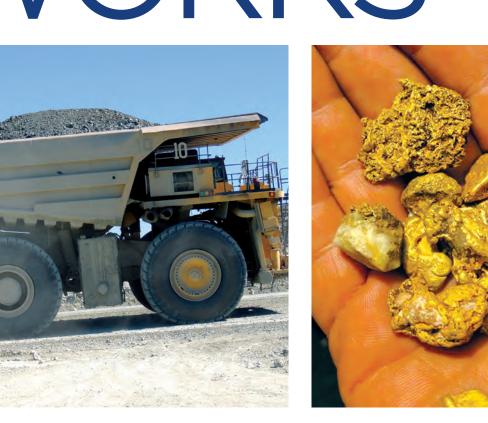
# PUBLISHED BY THE SOCIETY FOR MINING. METALLURGY & EXPLORATION

# **BY W. SCOTT DUNBAR**







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# Contents

Pr	eface
1	The Where, What, and How
	The Structure of Planet Earth1
	Earth History
	Definitions
	Concentrations of Metals
	Ore-Forming Processes
	Mine Exploration
2	Mining Methods
	Surface Mining
	Underground Mining
	Mining Operations and Equipment53
3	Mineral Processing
	and Metal Refining
	Comminution
	Flotation
	Smelting and Refining
	Hydrometallurgy
	Biohydrometallurgy
	Physical Separation Methods
4	Nonmetallic Minerals
	Coal
	Diamonds
	Aggregates
5	Mine Waste Management
	Mine Waste Streams and Disposal
	Mine Waste Treatment Methods

# iii

6	Mining, People, and the Environment	. 151
	Human Resources	. 151
	Safety	. 152
	Health	. 156
	Mining and Communities	. 158
	Sustainable Mining?	. 160
7	Mining and Money	.165
	Mineral Products	. 166
	Metal Prices	. 171
	Resources and Reserves	. 175
8	A Future of Mining	.185
	Drivers of Innovation in Mining	. 185
	Paths for Innovation	. 188
	Biotechnology to the Rescue?	. 191
	Can the Industry "Grow Its Own"?	. 195
	In Extremis	. 198
Ар	pendix A: All the Chemistry You Need to Know	. 201
	Atoms and lons	. 201
	Oxidation and Reduction.	. 203
	The Main Message	. 204
	Kitchen Chemistry	. 205
	Acid and Alkali Solutions	. 206
Ind	lex	. 207

iv

# Preface

Mining is interesting. It is a collection of processes that increase the very low concentrations of minerals and metals in the earth to levels that can be used in everyday life. In most cases, the increase in concentration is several orders of magnitude. The fact that it is physically, chemically, and economically feasible to do this is quite amazing. For some metals, the concentration process is like combing through a field of haystacks to find one needle every day for 10 years or more.

What is also interesting about mining is the different kinds of professionals required to make it happen: geologists, almost every kind of engineer or scientist, investment bankers, accountants, economists, health and safety specialists, community relations experts, and others. Furthermore, mining has economic effects at both global and local scales.

The intent of this book is to provide an understanding of the collection of processes, but to do so in a manner that makes the reader want to know more. The only requirement is an interest in science and technology. Very large and very small numbers are used to describe concepts. Some high school chemistry is used to describe processes, but it is not essential for understanding. Appendix A describes all the chemical concepts used.

The book consists of eight chapters. Chapter 1 provides an explanation of how mineral deposits are formed and how they are found. Chapter 2 describes mining methods, the systems and machines used to extract rock containing minerals of interest (called ore) from the earth. Chapter 3 describes a few of the methods used to process the ore and produce metals. Much of mining refers to the production of metals, but nonmetallic minerals are a huge part of the industry. Chapter 4 discusses the particular examples of coal, diamonds, and gravel (aggregates). In addition to minerals and metals, mining and processing ore produce large amounts of waste products that must be managed, often for an indefinite period. The science and technologies applied to this important part of mining are the topics of Chapter 5. Mining also involves people, the communities in which they live, and the government of the country in which the mine is located. As described in Chapter 6, the related issues are important, interesting, and sometimes challenging.

Chapter 7 describes the mechanisms for establishing prices of metals and minerals and discusses the legal definition of *ore*. Finally, Chapter 8 presents some ideas about mining in the future.

The majority of the images in the book were generously provided by many companies and individuals. Pictures in a book like this are worth far more than a thousand words, and the attributions under the figures do not begin to express my gratitude and appreciation. Many thanks are extended to those who helped illustrate the many topics. My colleagues Jocelyn Fraser, Mike Hitch, Bern Klein, Rick Lawrence, Allan Moss, and Dirk van Zyl helped with advice and ideas on content, but, of course, any errors are solely my responsibility.

I was encouraged to write this book by Jane Olivier, the books manager at SME, and I am so glad she asked. Diane Serafin edited the manuscript and kept me organized and her head cool despite delays and some software glitches.

The origins of this book lie with Steve Ralbovsky, a former partner at PricewaterhouseCoopers in Phoenix, Arizona, who in 1999 asked me to develop presentations that would explain mining to non-miners attending the PWC Annual School of Mines. For several years Steve twisted the dials and knobs to set the tone and level of the content. The presentations were mostly pictures, but over time I inserted many hidden slides with notes and explanations of the pictures. This was the starting point for this book. Thank you, Steve; it was fun.

In 2007, Simon Houlding of Edumine in Vancouver asked me to make a webcast from the slides, forcing me to learn to speak into a microphone and look into a camera at the same time, pretending that an audience was there. Simon then asked for an online course and quizzes. The information on the hidden slides helped, but good questions that don't frighten the learner are hard to compose. The staff at Edumine also encouraged me to keep the material up to date and to do a Spanish version of the webcast (No puedo. Si, se puede.) Challenges are good. Thank you Simon, Jennifer, Mariana, and Sandra.

Finally, thanks go to my wife Petra, who provided honest feedback, and to my sons and daughters for their comments, ideas, and inspiration.

## CHAPTER

# The Where, What, and How

Have you ever wondered why mines are located where they are?

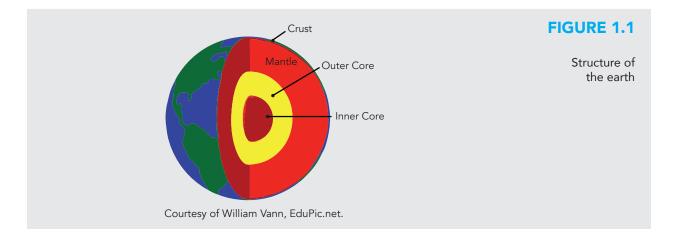
The ore deposits where mines are situated were formed millions of years ago in particular places by a number of different processes. Most of these processes continue to occur in the earth, but it can take millions of years for an ore deposit to form.

The material in this chapter will enable you to understand where and how ore deposits form and how geologists go about finding them. It will also define some common terms and concepts used in geology, mining, and mineral processing.

Let's start with what planet Earth looks like. It is the source of all the things we mine.

## THE STRUCTURE OF PLANET EARTH

The earth is composed of three concentric shells (Figure 1.1). The *crust* is the outer shell, 30–50 km thick under the continents and 5–10 km under the oceans. It is underlain by the *mantle* that extends from the crust to the interface between the mantle and the earth's core at a depth of 2,900 km. The earth's core, which has a radius of about 3,400 km, consists of an outer liquid core and an inner solid core. The radius of the earth is about 6,371 km.



The most interesting aspect of the geological processes involved in forming ore deposits is their scale, both in space and time. The formation of an ore deposit is essentially a concentration process that occurs within the crust and, in some cases, the upper mantle. However, these processes are fueled by heat generated deep within the earth's core and, in this sense, they are planetary in scale. The result is higher than average concentrations of minerals and metals at depths that are relatively shallow. The earth is a big planet. We have mined only as deep as 3–4 km, barely a pinprick.

## EARTH HISTORY

The earth is very old—about 4.6 billion years old. Table 1.1 shows Earth's history from the time it formed until the present. There are names for particular eons, eras, and periods of Earth's history. The origins of these names is often interesting. For example, the Ordovician period was named after a Celtic tribe called the Ordovices in order to resolve a dispute between geologists who were assigning ages to rock formations in northern Wales as either Cambrian, an earlier period, or Silurian, a later period.

The nickel deposits near Sudbury, Ontario, are believed to have been formed about 2 billion years ago or 2,000 MYBP (million years before present). Oil deposits in western North America formed between 360 and 400 MYBP, coal deposits formed

Eon	Era	Period	MYBP*	Major Events
	Cenozoic	Quaternary	0–1.6	Mastodons and humanoids
		Tertiary	1.6–66	Dinosaur extinction 66 MYBP
		Cretaceous	66–144	Ore deposits in western North America and South America
	Mesozoic	Jurassic	144–208	Coal deposits in British Columbia and western United States
		Triassic	208–245	Dinosaurs appeared 230 MYBP
Phanerozoic		Permian	245–286	
		Pennsylvanian	286–320	Coal deposits in eastern United States
		Mississippian	320–360	
	Paleozoic	Devonian	360–408	Oil deposits form in western North America
		Silurian	408–438	
		Ordovician	438–505	Trilobites appeared 500 MYBP
		Cambrian	505–543	
Proterozoic			543–2,500	Sudbury (Ontario) nickel deposits
Archean			2,500–4,600	

\*MYBP = million years before present.

Earth's history

between 150 and 200 MYBP, and several metallic ore deposits in North and South America were formed between 60 and 150 MYBP. Dinosaurs appeared about 230 MYBP and their extinction occurred about 66 MYBP, possibly due to the effects of a large meteor striking the earth. The evolution of humans is believed to have begun about 1.6 MYBP.

From Table 1.1 it is evident that geological processes are slow and span hundreds of millions of years. The processes of metal concentration are equally slow and might occur over a few hundred thousand to million years. An equally important concept is that the geological processes that occurred millions of years ago can also occur in the present. Hence the expression:

The present is the key to the past.

What this means is that the geological processes seen today likely occurred in the past. The climate might have been different, but the earth continued to do what it has always done. Consistency of geological processes over time is a useful concept in mine exploration.

## DEFINITIONS

Keep thinking about big spatial scale and very long times as you read through these definitions of terms that will be used often in this book.

## **Chemical Elements**

*Chemical elements* are a pure substance consisting of one type of atom. They are the building blocks of matter. Some examples are given here:

Al – aluminum	Mo – molybdenum
Ag – silver	O – oxygen
Au – gold	Pb – lead
Cu – copper	S – sulfur
Fe – iron	Si – silicon
H – hydrogen	Zn – zinc

Some of the symbols are simply the first letter of the name of the element. Others are derived from the Latin word for the element: **Arg**entum, **Au**rum, **Cu**prum, **Fe**rrum, **P**lum**b**um.

Figure 1.2 shows samples of particular elements. Native elements are those found in nature, but these are relatively rare and they are not pure. Elements of economic interest are often combined with other elements.

## FIGURE 1.2

Samples of copper, silicon, and gold



**Native copper nugget** about  $2.5 \times 3.5$  cm in size. Copper was the first element used by humans. It is combined with tin to form bronze and with zinc to form brass. Copper is the second best conductor of electricity, but it does corrode if exposed to moisture. (The best conductor is silver.)



**Pure silicon.** Silicon is one of the most abundant elements in the earth. It combines with oxygen to form quartz, a component of sand that is used to make glass. Since the mid-20th century silicon has been used in electronic devices.



**Pure gold leaf** about 1 cm long. Gold is very rare, which makes it valuable. It is also inert and malleable, which makes it useful for jewelry. Gold is the third best conductor of electricity behind silver and copper, but its inertness makes it useful in applications where corrosion would cause a problem.

Photos courtesy of Images-of-Elements.com 2014, CC BY 3.0.

# Compounds

Elements combine to form *compounds*. A compound contains at least two *different* elements. Some examples are given in Table 1.2. The numerical subscripts in the chemical formulas denote the number of atoms in the compound. If there is no subscript, then only one atom is present.

Compound	Chemical Formula
Salt: sodium (Na) and chlorine (Cl)	NaCl
Silver chloride: silver and chlorine	AgCl <sub>2</sub>
Vinegar (acetic acid): a bitter combination of carbon (C), oxygen, and hydrogen	CH <sub>3</sub> CO <sub>2</sub> H
Sulfuric acid: 2 hydrogen, 1 sulfur, and 4 oxygen atoms	$H_2SO_4$
Copper sulfate: 1 copper, 1 sulfur, and 4 oxygen atoms	CuSO <sub>4</sub>

## **Minerals**

The definition of a *mineral* is

A solid, naturally occurring, inorganic compound having a definite chemical composition (of elements) that may vary within limits.

# TABLE 1.2

Examples of compounds

## ARE YOU A BIT CONFUSED? WHAT IS A MOLECULE?

A molecule is formed when two or more atoms join together chemically. Compounds form when two or more *different* atoms join together. All compounds are molecules, but not all molecules are compounds. Molecular hydrogen (H<sub>2</sub>), molecular oxygen (O<sub>2</sub>), and molecular nitrogen (N<sub>2</sub>) are not compounds because each is composed of a single element. Water (H<sub>2</sub>O), carbon dioxide (CO<sub>2</sub>), and methane (CH<sub>4</sub>) are compounds because each is made from more than one element.

## Some examples of minerals are shown in Figure 1.3.



**Rose quartz**,  $SiO_2$ . The color, which may vary from pale pink to rose red, is due to trace amounts of metals such as iron, titanium, or aluminum. The color may be changed by irradiating the crystal with ionizing radiation. (Photo courtesy of Wikimedia Commons.)



**Pyrite**,  $FeS_2$ , is an iron sulfide sometimes known as "fool's gold" because it develops a sheen on its surface that looks like gold. The giveaway is the cubic crystals. Pyrite is the most common sulfide mineral. (Photo courtesy of Jonathan Zander 2007, CC BY-SA 3.0.)



**Chalcopyrite**,  $CuFeS_2$ . Except for the beak and eyes, this carving of an owl is made of chalcopyrite, a common copper sulfide found in most copper mines. The owl is about 6 cm high and is standing on quartz crystals. (Photo courtesy of Adrian Pingstone 2004, CC BY-SA 3.0.)

The following list describes other minerals of interest:

- Hematite (Fe<sub>2</sub>O<sub>3</sub>), an iron oxide, is the main source of iron. Hematite usually precipitates from standing water collecting in layers. It is typically dark colored, but weathering processes can cause it to become hydrated to form an iron hydroxide, essentially rust, that is responsible for the red color of many soils.
- Sphalerite (ZnS), which is zinc sulfide, and galena (PbS), which is lead sulfide, usually occur together in lead-zinc or lead-zinc-silver deposits.
- Pyrrhotite (Fe<sub>1-x</sub>S<sub>,</sub> where  $0 \le x \le 0.2$ ) is a magnetic iron sulfide with a strange chemical formula that seems to suggest partial atoms. However, this is merely

## FIGURE 1.3

Hand-size samples of minerals a convenient way to show a variable amount of iron. The end member of the series troilite FeS (x = 0) is nonmagnetic.

• The feldspar minerals—KAlSi<sub>3</sub>O<sub>8</sub>, NaAlSi<sub>3</sub>O<sub>8</sub>, CaAl<sub>2</sub>Si<sub>2</sub>O<sub>8</sub>—are silicates that comprise 60% of the earth's crust, a feature that distinguishes the crust from the mantle, which has fewer silicate minerals.

## **MORE ABOUT MINERALS**

The definition of *minerals* earlier in this chapter excludes substances formed by biogeochemical processes. However, there is a close link between the metabolic activities of microorganisms and mineral formation. In fact, microorganisms are capable of forming minerals and particular crystal structures that cannot be formed inorganically. Based on this, Skinner (2005) made the following expanded definition of a mineral:

An element or compound, amorphous or crystalline, formed through biogeochemical processes

where the prefix *bio* reflects the role of living systems in mineral formation. Skinner views all solids as potential minerals.

Minerals have been traditionally classified according to their chemical and crystal structure. For example, there are the silicates and the sulfides, the cubic and tetragonal crystal shapes, and so forth. This is a static classification system. Robert Hazen and John Ferry (2010) have shown that minerals evolved in a series of three eras beginning with an initial era of planetary accretion more than 4,550 MYBP when there were only about 60 minerals; through a second era of crust and mantle reworking to 2,500 MYBP, after which there were 1,500 minerals; and then into the current era in which living systems catalyze many mineral-forming reactions, resulting in the almost 4,400 mineral species known today. It is a very dynamic system that is ongoing.

#### Rock

A *rock* is a solid assemblage of minerals. A soil is also an assemblage of minerals but it is not solid. Soils are produced by weathering of rock, a part of the rock cycle that will be discussed later. There are many types of rocks, which will also be discussed later in context.

#### **Ore Deposit**

An *ore deposit* is a solid, naturally occurring mineral concentration useable as mined or from which one or more valuable constituents may be economically recovered.

The implications are that *current technology* and *economic conditions* make economic recovery possible. However, time is an important consideration. Technology can

turn an uneconomic deposit into an economic one. However, it can take some time for technology to develop into a practical form. Also, it is long-term economic conditions that govern whether a mine goes into operation and stays in operation. Short-term economic disruptions, such as the crisis in 2008, do not have much effect on mining operations.

This is not a legal definition of *ore*. The legal definition involves the issue of having the technical and legal ability to extract something of value from it. That aspect, and particularly the issue of how an ore deposit is reported to shareholders of a publicly traded mining company, are discussed in Chapter 7.

## **Base Metal and Precious Metal**

A *base metal* is a metal basic to industry and society such as iron, copper, lead, zinc, and aluminum. A *precious metal* is a rare, naturally occurring metallic element of high economic value such as gold and silver and the platinum group metals: palladium, rhodium, iridium, osmium, and ruthenium.

## What Is Ore Grade?

*Ore grade* is the concentration of economic mineral or metal in an ore deposit. A fundamental unit used in defining grade is tonnage. The definition of *one ton* has changed over the years but can be summarized as follows:

- One US ton (1 ton) is 2,000 lb; also called a short ton.
- One UK or Imperial ton is 2,240 lb; also called a long ton—obsolete since about 1965 when the United Kingdom adopted the metric system.
- One metric ton (1 t) is 1,000 kg. The term *tonne* is still used in many countries.

Given that 1 kg is about 2.2 lb, 1,000 kg or 1 t is about 2,200 lb, or about 1.1 times the mass of 1 US ton. (It is close to the mass of the Imperial ton, but there's no need to be confused with an obsolete unit.)

The grade of a base metal deposit is expressed as a mass percentage. For example, a fairly rich copper deposit might have a grade of 0.4%. This means there are 0.004  $\times$  1,000 kg = 4 kg of copper in each metric ton of ore. Alternatively there are 0.004  $\times$  2,000 = 8 lb of copper in each US ton of ore. This all works out as shown in the following equation—watch the units (remembering that "t" is the metric ton and "ton" refers to the US ton):

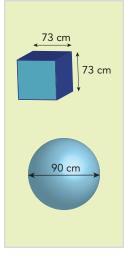
$$4 \text{kg/t} = \frac{4 \text{ kg} \times 2.2 \text{ lb/kg}}{\text{t}} = \frac{8.8 \text{ lb}}{1 \text{ t} \times 1.1 \text{ ton/t}} = 8 \text{ lb/ton}$$

#### **TROY OUNCES**

The origin of the word *troy* comes from the city of Troyes in France, an important trading city in the Middle Ages. A troy ounce is about 10% larger than the avoirdupois ounce used to measure the weight of things like the weight of the human body or food items. However, if you consider your body precious, start measuring its weight in troy ounces, convert those troy ounces to troy pounds, where 12 troy ounces equals one troy pound, and you can fool yourself into thinking that dieting is unnecessary.

The grade of a precious metal deposit is expressed in grams per metric ton or troy ounces per ton, where the common abbreviations are g/t and oz/ton, respectively. The ounces are troy ounces where 1 oz = 1.097 ounces avoirdupois = 31.1 g. For example, the gold grade at the Eskay Creek mine in British Columbia was about 0.96 oz/ton or 29.856 g/t, which is a very rich ore. A more typical gold grade is 1-4 g/t for open pit mines, 4-6 g/t for marginal underground mines, and 8-10 g/t for high-quality underground mines.

Because there are 1,000 kg in a metric ton and 1,000 g in a kilogram, there are 1 million grams in a metric ton. Thus, if gold grade is given in grams per metric ton, then it is in parts per million (e.g., 6 g/t is 6 ppm).



#### HOW BIG IS 1 METRIC TON OF ROCK?

One cubic meter (m<sup>3</sup>) of solid rock weighs between 2.5 and 3 t. This is the density of rock. Assume the density is 2.6 metric tons per cubic meter or 2.6 t/m<sup>3</sup>. Invert this to get 1/2.6 =  $0.385 \text{ m}^3$ /t. This means that 1 t of rock occupies a volume of  $0.385 \text{ m}^3$ . This volume could be any shape. If it's a cube, then one side of the cube is  $0.385^{1/3} = 0.73 \text{ m}$  or 73 cm. Such a cube is shown to the left.

If the volume is a sphere, then because the volume of a sphere of radius r is given by  $4\pi r^3/3$ , the radius of the sphere can be computed as  $(3 \times 0.385/4\pi)^{1/3} = 0.45$  m or 45 cm. Its diameter is 90 cm. Such a sphere is shown to the left.

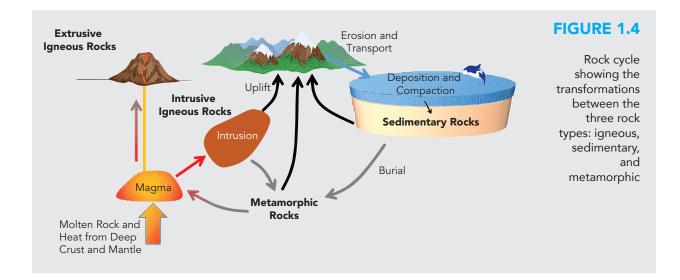
You can extend a measuring tape to these dimensions to get an idea of how big these volumes are.

To convert a weight percentage to parts per million, multiply the percentage by 10,000, that is, 10,000 ppm = 1%. Thus 0.4% is  $0.4 \times 10,000 = 4,000$  ppm. Conversion of base metal grades to parts per million is not commonly done, but it serves to show the large difference—three orders of magnitude—between typical gold grades and typical base metal grades.

Diamond mine grades are measured in carats per metric ton (carat/t). Since 1 carat = 0.2 g, a diamond grade of 4 carats/t =  $4 \times 0.2 = 0.8$  g/t or 0.8 ppm.

## Rocks and the Rock Cycle

There are many physical and chemical cycles that occur in the earth (e.g., hydrologic cycle, carbon cycle). A rock cycle is illustrated in Figure 1.4.



Three types of rock are shown in the rock cycle: igneous, sedimentary, and metamorphic. *Igneous rocks* are formed when rock melts at temperatures between 600 and 1,200°C. The melt is called *magma* if it stays within the earth; if it exits the earth, as in a volcano eruption, it is called *lava*. Melted rock is light and will flow toward the surface where, if it solidifies within the earth, it is called an *intrusive igneous rock* (or simply an intrusion). If it exits the earth, it solidifies very quickly and is called an *extrusive igneous rock*.

Figure 1.5 shows some examples of igneous rocks. The magma that formed the granite cooled slowly, allowing the large crystals to form. Figure 1.5 also shows rhyolite, which formed by fast cooling of lava from a volcano, resulting in small crystals. Both rocks have about the same chemical composition. Figure 1.5 includes a sample of basalt, a volcanic rock that contains less than 50% silica and a lot of heavy dark minerals that contain iron and magnesium. While granite is the most common rock in the continental crust, basalt is the most common rock in the continental crust, basalt is the most common rock in the oceanic crust. Figure 1.5 also shows basaltic lava that, because of its low silica content, flows very easily. The Columbia plateau in southern Washington state and adjacent Oregon and Idaho is a large (163,000 km<sup>2</sup>) "flood basalt" formed during volcanic

## FIGURE 1.5

Examples of igneous rocks





Granite

**Rhyolite (likely hand-size specimen)** (Courtesy of US Geological Survey.)



Basalt (Courtesy of US Geological Survey.)

Basaltic lava (Courtesy of US Geological Survey.)

activity 15 to 17 million years ago. Hawaii was formed and continues to be formed by basalt flows.

*Sedimentary rocks* consist of particles formed by weathering of rock transported by flowing water and deposited either on the earth's surface or at the bottom of a body of water. Unless disrupted by some other process, the deposition results in layers of sediments called *strata*. As more sediments are deposited, the layers become compacted and hard. An example of a sedimentary rock formation is shown in Figure 1.6. The sediments forming these rocks were deposited over several million years during the Triassic period, about 200–245 MYBP.

The appearance of the sediments reveals a lot about the environment in which the sediments were deposited. The lower red-colored rocks in Figure 1.6 are composed of silt (siltstones) and were deposited in a shallow, calm marine environment. The red color is due to iron minerals (hematite) formed by exposure to the atmosphere, essentially a rusting process. The upper gray-colored rocks are limestone composed



Courtesy of Mark A. Wilson via Wikipedia.

FIGURE 1.6

Part of the Virgin Formation, a sequence of sedimentary rocks in southwestern Utah formed during the Triassic period, about 200–245 MYBP

of the skeletal remains of marine organisms. These sediments were deposited in deeper water where the iron present was not exposed to the atmosphere.

Over time, igneous and sedimentary rocks may become buried and subjected to high pressures and temperatures, which cause physical and chemical transformations but do not melt the rock. The result is a *metamorphic rock*. The variety of sources combined with the large possible range of temperatures and pressures gives rise to a wide variety of chemical and crystal textures of metamorphic rocks. Two examples are shown in Figure 1.7.



**Quartzite**—originally sandstone (Photo courtesy of Siim Sepp 2005, CC BY-SA 3.0.)



**Gneiss**—originally igneous or sedimentary. The bands are caused by high shear forces that cause the rock to behave like plastic. (Photo courtesy of Siim Sepp 2005, CC BY-SA 3.0.)

Metamorphic rocks comprise about 85% of the earth's crust, and it is estimated that about 10% of the earth's crust is composed of sedimentary rocks. These sedimentary rocks must exist at or near the surface of the earth because at depth they would be exposed to high pressure and temperatures and become metamorphosed or melt to become igneous rocks.

FIGURE 1.7

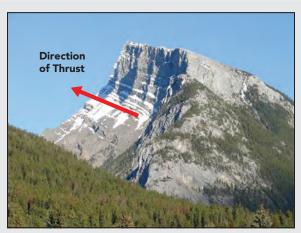
Two examples of metamorphic rocks As may be seen in Figure 1.4, the rock cycle actually consists of a number of subcycles. Rocks are continually transforming from one type to another over millions of years. Although the source of heat is deep within the earth (due to heat generated by radioactive decay), most of the processes shown in the rock cycle occur at depths less than 20 km.

## Faults, Fractures, and Folds

The earth's crust is subjected to very large forces that lead to ruptures called *faults* and numerous *fractures*. There are several well-known faults, such as the San Andreas fault in California where two large plates of the earth's crust slide past one another and cause earthquakes. The Rocky Mountains in Canada were formed by thrusting of sedimentary rocks from the west to the east. Mount Rundle, one mountain in the Canadian Rockies, is shown in Figure 1.8. The uplifted layers can be seen on the left, the east side. The layers have been thrust from west to east along a subhorizontal fault over existing rocks, a process that has resulted in older sedimentary rocks being placed on top of younger ones, the reverse of what happens when sedimentary rocks are formed.

## FIGURE 1.8

View looking south of Mount Rundle (Banff, Alberta), formed by a series of thrust faults



Adapted from Antony Stanley 2006, CC BY-SA 2.0.

Fractures ("mini-faults") also result from the large forces imposed on rocks. Fluids with mineral solutions percolate through these fractures and deposit minerals or metals in the fractures, forming veins. An example of small veins containing visible gold is shown in Figure 1.9. (Visible gold is actually a rare occurrence.)

Stresses combined with temperatures and pore water pressure can soften rocks so that they bend easily or actually flow. Layers of sedimentary rocks that were originally horizontal can become *folded*, as shown in Figure 1.10. Folds are easily seen



Courtesy of Pretium Resources.

# <image>

Courtesy of Alberta Geological Survey.

in sedimentary rocks, but they also occur in metamorphic rocks and also form in igneous rocks subjected to stresses in the later stages of magma cooling.

## **CONCENTRATIONS OF METALS**

Geochemical models can be used to estimate the average concentrations of metals in the crust of the earth. The details of these models are complex. However, the main observation is that the typical grade of an ore deposit is larger, sometimes much larger, than the average concentration of the metal in the earth's crust.

## FIGURE 1.9

Gold veins at Brucejack property, northwestern British Columbia

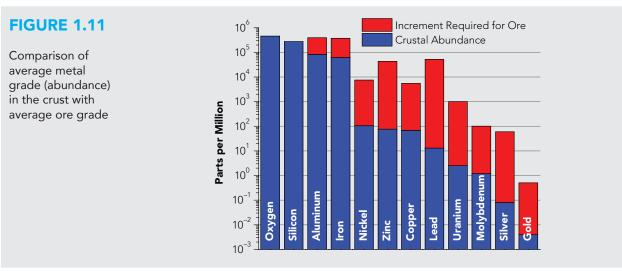
# **FIGURE 1.10**

Folded sedimentary rocks at Mount Head in the Canadian Rockies For example,

The grade of a typical gold deposit is of the order of thousands of times greater than the average concentration of gold in the earth's crust.

In contrast, the grade of an iron deposit is only 10 times greater than the average concentration of iron in the earth's crust.

These relationships are illustrated for different elements in Figure 1.11. Note the logarithmic scale where one major division denotes an order-of-magnitude difference (e.g., from  $10^1 = 10$  to  $10^2 = 100$ ). This scale is used because of the large range of concentrations.



The abundances of oxygen and silicon are shown for comparison:

Silicon ~300,000 ppm, Oxygen ~500,000 ppm

These two elements occur as components of minerals; that is, they are bound chemically to other elements.

The concentrations of metals in ore deposits depend on the current price of the metal because ore is what can be mined under current economic conditions. Thus, as the price goes up, the concentration of metal required to make rock an ore deposit goes down. However, relative to the concentrations in the earth's crust, the changes are not significant.

# Index

Note: f. indicates figure; t. indicates table

## Α

acid, in rivers, 142 acid rock drainage (ARD), 141-143, 143f. acid solutions, 206 acquisition, of mining companies, 26 activated carbon, 98-102 aerobic processes, of waste treatment, 149 aggregates business of, 125-126 origin and occurrence, 124 processing, 125, 125f. uses for, 124, 124t. alkali solutions, 206 anaerobic processes, of waste treatment, 149 anomalies, 29, 30f. aquifers, 62 ash, 114 asthenosphere, 18 at-mine return (AMR), 168 atomic weight, 81 atoms, 201-203, 201f. aurocyanide and carbon adsorption, 98-102 as by-product of gold ore processing, 96-98 Australian Academy of Science, 187 automation, of the mining cycle, 189, 190f.

## В

bacteria *Delftia acidovorans*, 193 methanogens, 194 oxidation by, 103–106 *Ralstonia metallidurans*, 192, 192*f*., 193*f*.

sulfate-reducing bacteria (SRB), 147-148, 147*f*., 193 basalt, 9, 10f. basaltic lava, 9-10, 10f. base metals defined, 7 separation of from precious metals, 83 benches, pit slope, 39, 39f., 40f. biohydrometallurgy, 103-106 biomining, 105, 106f. "bio-ore," 194 biotechnology, 191-196, 197f. black smokers, 21 blasting, 55-56, 55f., 56f., 65f. blister copper, 85 block caving, 50-52, 51f., 52t. block models, 34-35, 35f., 180f. Blusson, Stewart, on finding diamonds in the Canadian Arctic, 121 British Columbia, Canada, copper-gold occurrences and mines map, 27f. by-products, 169-170

# С

Canadian Securities Administrators, 181 capital productivity, 188–190 carbon adsorption vs. absorption, 100 carbon in pulp (CIP) process, 99, 100–101, 101*f.* carbon-in-column method, 99–100 carbon-in-leach (CIL) process, 99, 100 processes, 98–99, 99*f.* cave front, 51 centrifugal concentrators, 108, 108f. chalcopyrite, 5f. chemical elements defined, 2 examples of, 4f. Chicago Mercantile Exchange (CME) Group, 174-175 classification (separation process), 74 clays, 113 cleaner cells, 79 coal formation of, 113-115, 114f. industry characteristics, 115 processing, 115-119, 116f. rank, 113, 115f. as rock vs. mineral, 116 run-of-mine (ROM), 116-117 washing, 116 Code for Reporting of Mineral Resources and Ore Reserves (JORC), 183 collectors (flotation chemical), 77, 78, 79 comminution crushing, 69–71, 122–123 cyclones (hydrocyclones), 74-75, 75f. defined, 67 grinding, 71-74 grinding circuits, 75-76, 76f. physics of, 72 commodities, 165, 165*f*. commodity exchanges, 171, 174-175 communities health in, 156–158 and mining, 158-160 compounds defined, 4 examples of, 4t. concentrates, 78, 78f. See also copper concentrates; sulfide concentrates concentration calculating concentration factor (CF), 80-81, 80f. relationship to ore grade, 13-15, 14f. concentrators, 68f., 108-109, 108f., 109f.

contact metamorphism, 24 continental boundaries, 22, 22f. Cooney, James, on social license to operate (SLO), 158 copper example of, 4f. plating experiments, 205-206, 205f. prices, 175, 175f. copper concentrates bio-heap leaching, 106 refining, 86-88, 88f. smelting, 85-86, 86f. solvent extraction and electrowinning (SX/ EW), 91-94, 92f., 93f., 95f. co-products, 169-170 core, earth, 1 corporate social responsibility (CSR), 156, 159 costs, reduction of, 189 cratons, 119 crosscuts, 50 crushers cone, 70–71, 70f. gyratory, 69-70, 69f. high-pressure grinding rolls (HPGR), 122-123, 122f. jaw, 70, 70f. crust, earth, 1, 18-22 cut-and-fill mining, 46, 47f. cutoff grade, 179-180, 180f. cyclones (hydrocyclones), 74-75, 75f.

# D

dams, tailings centerline construction, 134–135, 134*f*. downstream construction, 134, 134*f*. history of, 136 upstream construction, 133–134, 133*f*. *Delftia acidovorans*, 193 delivery contracts, 167 dense media separation (DMS), 118, 118*f*. developing countries, and mining, 157–160 dewatering, 61–62, 63*f*. diamonds age-dating, 120 Canadian Arctic deposits, 120–121 formation of, 119-120, 120f. processing, 122-124 dilution, 62, 64, 64f., 65f. dirigibles, 65, 66*f*. disclosure, 181-183 disseminated deposits, 17 Dobbs, Richard, on resource curse, 160 dose, defined, 157 drilling cost considerations, 33-34 diamond core, 31, 32f., 33 equipment, 32f., 54-55, 54f., 55f. goal of, 34 reverse circulation, 33, 33f. dry stack, tailings, 137, 138f. Dupont equation, 188

## Ε

earth history of, 2-3, 2t. plate tectonics, 18-22 structure of, 1-2, 1f. economic analysis, 182-183, 182f. economic conditions, effect of on mining operations, 6-7 electrolysis, 204 electrons, 201-203, 201f., 202f. electrorefining, as different from electrowinning, 92-93 electrowinning, 89, 94, 101-102, 101f., 166-167 See also solvent extraction and electrowinning (SX/EW) elements, 2, 4f. Elsner reaction, 90 emergence, 196 employees health of, 156 and safety behavior, 154-155 engineering, and pit slope stability, 39-40 engineers, 152

environmental engineers, 129 epithermal deposits, 17 Equator Principles, 158 equipment autonomous haul trucks, 189, 190f. centrifugal concentrators, 108, 108f. concentrators, 68f. crushers, 69-71, 69f., 70f. cyclones (hydrocyclones), 74-75, 75f. dense media separation (DMS) machines, 118, 118*f*. drills, 32f., 54-55, 54f., 55f. filter press, 137, 138f. high-pressure grinding rolls (HPGR), 122-123, 122f. high-tensile steel mesh, 60, 60f. load-haul-dump (LHD) machines, 47, 58-59, 58f., 59f., 60, 65-66 magnetic separators, 109-110, 110f. oxygen flash furnace, 85f. rock bolts, 61, 61f. rotary breakers, 117, 117f. screens, 117, 118f. shaking tables, 106-107, 107f. shearers, 50, 50f. sluice boxes, 107, 107*f*. spiral concentrators, 108-109, 109f. X-ray separators, 123–124, 123f. evaporite deposits, 24 experiments, copper plating, 205–206, 205f. exploration drilling, 31, 32f., 33–35, 33f., 34f. expenditures vs. discoveries, 186-187, 186f. geochemical surveys, 31, 31f. geological mapping, 28–29 geophysical surveys, 29-30, 29f., 30f. and reserve ratios, 179 role of in mining, 26–28 satellite imaging, 28, 28f. See also reserves; resources external dilution, 64, 64f. extrusive igneous rocks, 9

# F

faults, 12, 12f. feasibility studies, 181-182 feldspars, 6, 17, 113 Ferry, John, and mineral classification, 6 filter cakes, tailings, 137 filters, tailings, 131-132 Fipke, Charles, on finding diamonds in the Canadian Arctic, 121 Fire into Ice: Charles Fipke and the Great Diamond Hunt (Frolick), 121 flocculants, 110 flotation cells, 79, 80f. circuits, 78-81, 79f. of coal, 118–119 and grade recovery, 84 mineral separation in, 81-83 process, 77-78, 77f. folds, 12-13, 13f. fractures, 12 freeboard, 132f., 133 Frolick, Vernon (Fire into Ice), 121 froth flotation. See flotation frother, 77

# G

gangue, 67, 69f. geochemical surveys, 30, 31f. geological mapping, 28-29 geological processes, slowness of, 2 geophysical surveys, 29-30, 29f., 30f. geotechnical engineers, 129 glacial till, 121, 121f. gneiss, 11f. gold alternatives to cyanide leaching, 102-103 aurocyanide and carbon adsorption, 98-102 and bacterial growth, 192 biooxidation, 105 carbon stripping and electrowinning, 101–102, 101*f*., 102*f*. cyanide leaching, 96-97 example of, 4f.

Merrill-Crowe process, 97–98, 97*f*. prices, 173–174, 173*f*. smelting, 102, 103*f*. and termite mounds, 193 gossan deposit, 26 Gowan, Mike, on co-disposal, 146 grade. *See* ore grade granite, 9, 10*f*. gravel pits, 126 greenhouse gas (GHG) emissions, 161, 161*f*. grinding autogenous (AG), 71–73 ball mill, 73–74, 74*f*. circuits, 75–76, 76*f*. semiautogenous (SAG), 71–73, 73*f*.

## Η

hazard, defined, 153 Hazen, Robert, and mineral classification, 6 health community, 156-158 occupational, 156 personal, 156 See also safety Horn, Robert, on exploration challenges, 35 human resources, 151-152 hydraulic fill, 46 hydrometallurgy defined, 89-90 leaching, 90-91 hydrothermal metamorphism, 24 hydrothermal systems, 17 hydrothermal vents, 20f., 21 hyperaccumulators, 194-195

## I

igneous rocks defined, 9–10 examples of, 10*f.* igneous-hydrothermal ore deposits, 16–17, 16*f.* indicator minerals, 121 industrial minerals, 113 *Industry Guide 7* (U.S. Securities & Exchange Commission), 183

## Index

inertinite, 115 insect swarms, as model for biotechnology, 196 intergenerational equity, 161 internal dilution, 64, 64*f*. internal erosion, 132 intrusive igneous rocks, 9 ions, 201–203, 202*f*., 203*f*. "iron hat" deposits, 26 island arcs, 18*f*., 19, 22, 22*f*.

## J

Joint Ore Reserves Committee of the Australasian Institute of Mining and Metallurgy (JORC), 176, 183 joint products, 169–170 JORC diagram, 176

# Κ

Kessler, Stephen, on deposit discovery rates, 187 "Kilobots," 196, 197*f.* kimberlite pipes, 119–120 *See also* diamonds

## L

laterites, 26 lava, 9 leachate, 91 leaching alternatives to cyanide, 102-103 bio-heap, 106 biohydrometallurgy, 103-106 carbon adsorption processes, 98-102 cyanide, 96-103 Merrill-Crowe process, 97–98, 97f. in pads, 90-91, 91f. pressure, 94-96 reactions, 90 in-situ, 90-91 solvent extraction and electrowinning (SX/EW), 91-94 lifts, 90 liptinite, 115 lithosphere, 18

lixiviant, 91 London Bullion Market Association (LBMA), 172–173 London Metal Exchange (LME), 174–175 London Platinum and Palladium Market (LPPM), 172–173 longhole stoping, 48 longwall mining, 49–50, 50*f*:

## Μ

macerals, 114 magma defined, 9 and ore deposit formation, 16-17, 16f. magnetic separators, 109-110, 110f. management, 152 mantle, earth, 1, 119 massive sulfide deposits formation of, 17 volcanic (VMS), 21, 21t. matte, 85 metal equivalents, 170-171 metal streaming, 183 metals calculating deposit grade, 7-9 concentration of, 13-15 defined, 7 and microorganisms, 191-194 noble, 204 and plants, 194-195, 195f. prices, 171-175 refining, 86-89 separation of, 83 solubility of, 204 metamorphic rocks defined, 11 examples of, 11f. metamorphic deposits, 24, 25f. methanogens, 194 microorganisms, 191-194 mid-ocean ridges, 18f., 19-21, 20f. mine planning, use of block models in, 34-35, 35f., 180f. mine waste acid rock drainage (ARD), 141-143 calculating how many tons, 141

co-disposal, 146-147, 146f. health effects of, 157 management of, 129 passive treatment systems, 148-149 pit disposal, 145 reclamation, 144-145 sequestration, 145-146 streams of, 130 tailings, 130-139 treatment methods, 144-149 waste rock, 139-141 wastewater, 141, 147-149 mineral products by-products, 169-170 co-products, 169-170 examples of, 165f. and metal equivalents, 170-171 product streams, 166-167 and refinery contracts, 167-168, 168f. and smelter contracts, 166, 167-168, 167*f*. supply system, 166f. minerals calculating deposit grade, 7-9 classification of, 6 defined, 4 deposit discovery rates, 186-187 descriptions of, 5-6 examples of, 5f. indicator, 121 industrial, 113 separation of by flotation, 81-83 separation of by physical methods, 106-111 Skinner's expanded definition, 6, 116 See also mineral products mining cycle blasting, 55-56 drilling, 54-55 load-haul-dump (LHD), 58-59, 65-66 stages of, 53-54, 53f. support installation, 59-61 ventilation and refrigeration, 56-57 mining industry asset reduction in, 190 and biotechnology, 191-196

cost reduction in, 189 deposit discovery rates, 186–187 future of, 197*f*., 198 innovation in, 185 paths for innovation in, 188–190 and sustainability, 160–161, 186 Minto Deposit, resource classification example, 177, 178*f*. "modifying factors," 176, 176*f*. molecules, defined, 5 molybdenite, 82 muck, 48

## Ν

National Instrument 43-101 (NI 43-101), 181, 183 native elements, 2 net smelter return (NSR), 168 noble metals, 204 nonmetallic minerals, 113 North America, subduction zones in, 23*f*.

# 0

occupational health, 156 open pit mining pit slopes, 38-40, 38f., 39f., 40f., 41 pushbacks, 40-41, 40f., 42f. strip ratios, 42-43, 42t. typical configuration, 37-38, 37f., 38f. operating productivity, 188-190 operations blasting, 55-56 dewatering, 61-62 drilling, 54-55 grade control, 62, 64 load-haul-dump (LHD), 58-59, 65-66 support installation, 59-61, 63 ventilation and refrigeration, 56-57 ore, legal definition, 7, 175 ore deposits correlation with volcanic activity, 15, 15f., 16f. defined, 6-7 formation of, 2, 15, 19 igneous-hydrothermal, 16-17, 16f. massive sulfide, 17

## Index

metamorphic, 24, 25f. oxide, 25-26, 25f. sedimentary, 23-24 See also concentration; ore grade ore grade calculating, 7-9 control, 62, 64 cutoff grade, 179-180, 180f. defined, 7 flotation recovery and, 84 relationship to concentration, 13-15, 14f. ore-body models, 34-35, 34f. overflow, 75 oxidation, 25-26, 25f., 203-204, 203t. by bacteria, 103–106 oxygen flash furnace, 85f.

## Ρ

panel caving, 52 panels, 49, 52 paradox of plenty, 159-160 partial melt, 17 paste, tailings, 137, 137f. paste backfill, 46 percentage paid, 168 pH scale, 206, 206t. phytomining, 194-195, 195f. phytoremediation, 194–195, 195f. pillars, 48, 49, 49f. pit slopes, 38-40, 38f., 39f., 40f., 41 placer deposits, 23, 24f. plants, metals in, 194-195 plate tectonics, 19 plates boundary types, 18f., 19-22 movement of, 18-19 precious metals defined, 7 separation of from base metals, 83 preconcentration, 189-190 prefeasibility studies, 181 pregnant leach solution, 91 prices/pricing commodity exchange, 171, 174-175 copper, 175, 175f.

gold, 173–174, 173*f*. independent, 171, 172–174 negotiated, 171 producer, 171 processing aggregates, 125 coal, 115–119 comminution, 67–76 diamond ore, 122–124 flotation, 77–84 gold ore, 96–103 guiding principle, 76, 79, 102 production level, 50 pushbacks, 40–41, 42*f*. pyrite, 5*f*, 82, 83*f*.

## Q

qualified person (QP), 181 quarries, 126 quartz, rose, 5*f.* quartzite, 11*f.* 

## R

Ralstonia metallidurans, 192, 192f., 193f. reclamation, 144-145, 144f. reduction, 203 refinery contracts, 167, 168-169, 168f refining of copper concentrates, 86-88, 88f. by electrolysis, 204 of lead concentrates, 88 of lead-zinc concentrates, 89 of nickel sulfide concentrates, 89 of zinc concentrates, 88-89 refrigeration, 57 reporting standards, 181-183 reserve ratios, 178-179, 179t. reserves classification of, 176-177 defined, 175-176 reporting standards, 181-183 See also exploration residence time, 76, 80 resource curse, 159-160

resources classification of, 176-177 defined, 175 in difficult places, 188, 188f. Minto Deposit classification example, 177, 178f. reporting standards, 181-183 See also exploration rhyolite, 9, 10f. Richards, J. P., on deposit discovery rates, 187 rift zones, 18f., 19 risk, 153 risk management, 153-154, 153f. rock bolts, 61, 61f. rock cycle, 9-12, 9f. rock(s), defined, 6 room-and-pillar mining, 49, 49f. rose quartz, 5f. rougher cells, 79

## S

safety, 152-155 See also health safety culture, 154-155 SAMREC Code, 183 sand and gravel. See aggregates satellite imaging, 28, 28f. scavenger cells, 79 sedimentary rocks defined, 10-11 examples of, 11f. sedimentary ore deposits, 23-24 sequestration, of acid-generating waste, 145–146, 146*f*. shaking tables, 106-107, 107f. shared value opportunities (SVO), 159 shields, 49 shotcrete, 61, 62f. shrinkage stoping, 48 silicon, 4f. Skinner, H.C.W., expanded definition of minerals by, 6, 116 slag, 85, 86 sluice boxes, 107, 107f. smelter contracts, 166, 167–168, 167*f*.

smelting chemical reactions in copper smelting, 87 of copper concentrates, 85-86, 86f. gold, 102, 103f. of lead concentrates, 88 of lead-zinc concentrates, 89 of nickel sulfide concentrates, 89 of zinc concentrates, 88-89 soap, 79 social license to operate (SLO), 158-159 soil(s), defined, 6 solvent extraction and electrowinning (SX/ EW), 91–94, 92f., 93f., 95f. spalling, 60 spiral concentrators, 108-109, 109f. split-off point, 169 stockpiles, 71, 71f. stoping, 48, 48f. strata, 10 Stratum International, mining professional survey, 152 strip ratios, 42-43, 42t. subduction zones, 18f., 22, 22f., 23f. sulfate-reducing bacteria (SRB), 147-148, 147f., 193 sulfide concentrates biohydrometallurgy, 103-106, 105f. mineral recovery from slurry, 84, 84f. pressure leaching, 94–96, 95f. smelting, 84-86 use of flotation to separate, 77-78 sulfuric acid, 85-86 surface mining. See open pit mining sustainable development, 160-161 sustainable mining, 160-161, 186 SX/EW. See solvent extraction and electrowinning (SX/EW)

# Т

tailings dams, 131, 133–135, 133*f*., 134*f*., 136 defined, 38, 129, 130 example of, 129*f*. forms of, 131, 131*t*. impoundment, 131, 132*f*.

## Index

impoundment water balance, 135, 135f. pit disposal, 145, 145f. pond, 133 thickened, paste, and filtered disposal, 135, 137, 137f., 138f., 139 wet disposal, 131-133, 132f. technical reports, 181-183 termites, and gold, 193 thickeners and clarifiers, 110-111, 110f., 111*f*. timber supports, 63 ton/tonne defined, 7 size of one metric ton, 8 toxicity, 157 treatment and refining charges (TC/RC), 168 troy ounces, 8

## U

undercut, 50 underflow, 75 underground mining block caving, 50–52, 51*f*., 52*t*. comparison of methods, 52–53, 53*t*. cut-and-fill, 46, 47*f*. and exploration, 44–45 longwall, 49–50, 50*f*. narrow vein mining, 47, 47f. resources, 52 room-and-pillar, 49, 49*f*. shaft construction, 45–46, 46*f*. stoping, 48, 48*f*. support installation, 59–61, 63 terminology, 44*t*., 45*f*. typical configuration, 43–44, 43*f*.

# V

VALMIN Code, 183 valuation, of by-products and co-products, 169–170, 170*f.* veins, 13*f.* ventilation, 56–57, 57*f.* vertical crater retreat (VCR), 48 vitrinite, 114–115, 115*f.* volcanic massive sulfide (VMS) deposits, 21, 21*t.* volcanos map of active, 15*f.* and ore deposit formation, 15–17, 16*f.* 

## W

waste rock/waste dump, 37, 139, 140*f*., 141 pit disposal, 145, 145*f*. *See also* strip ratios wastewater, 141, 147–149 wetlands, 148–149, 148*f*. workforce, 151–152

# Х

X-ray fluorescence, 123-124, 123f.

# Ζ

zinc, 98